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Global chronostratigraphical correlation table for the last 2.7 million years

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The table provides a correlation of chronostratigraphical subdivisions of late Cenozoic geological time, spanning the last 2.7 million years. The formal division of the Quaternary is the responsibility of the International Commission on Stratigraphy's (ICS) Subcommittee on Quaternary Stratigraphy (SQS), in partnership with the International Union for Quaternary Research's (INQUA) Commission on Stratigraphy and Chronology (SACCOM). This is the third published version of the chart. Earlier versions are Gibbard et al., 2004, 2005. See <http://www.quaternary.stratigraphy.org.uk/correlation/> for history.

Chronostratigraphy and the base of the Quaternary

The timescale is based on the internationally-recognised formal chronostratigraphical/geochronological subdivisions of time: the Phanerozoic Eon/era; the Cenozoic Era/epoch; the Quaternary System/Period; the Pleistocene and Holocene Series/Epoch, and finally the Early/Lower, Middle, Late/Upper Pleistocene Subseries/Subepoch. At present the Subseries (Subepoch) divisions of the Pleistocene are being formalised. Series, and thereby systems, are formally defined based on Global Stratotype Section and Points (GSSP) of which two (Holocene and Pleistocene Series) have been ratified for the Quaternary System. The currently ratified base of the Pleistocene is defined in a GSSP at Vrica in Southern Italy (Aguirre and Pasini, 1985), with an age of ~1.8 Ma. However, it is proposed to define the base of the Quaternary/Pleistocene at 2.58 million years from Monte San Nicola, also in southern Italy, which is the current GSSP for the Pliocene Gelasian Stage (Rio et al., 1994, 1998).

Since 1948 there has been a consensus that the boundary should be placed at the first evidence of climatic cooling of ice-age magnitude. This was the original basis for placing the boundary at ~1.8 Ma in marine sediments at Vrica in Calabria, in Italy (Aguirre & Pasini, 1985). It is now known that a major cooling occurred earlier, at c. 2.55 million years in the Mediterranean (Cita, this issue), and even earlier cooling events are known from the Pliocene. Since its definition at 1.8 Ma there has been strong pressure for the basal Quaternary/Pleistocene boundary to be moved downwards to better reflect the initiation of major global cooling (Pillans and Naish, 2004; Gibbard et al., 2005; Bowen & Gibbard, 2007), effectively corresponding to the Gauss/Matuyama magnetic Chron boundary (e.g., Partridge, 1997; Suc et al., 1997). See discussion in this volume (Ogg & Pillans, Head et al., Lourens).

At the time of publication discussions concerning the position of the boundary(ies) are at an advanced stage. The formal boundary of the Quaternary/Pleistocene at 1.8 million years is indicated but the proposal to define formally the base of the Quaternary from that of

the Pleistocene, at the 2.6 million-year boundary, has already been accepted following formal voting by the ICS. The proposal is awaiting formal ratification by the IUGS. It is expected a decision will be made following the International Geological Congress to be held in Oslo in August 2008.

Similarly there is uncertainty about the status of the term Tertiary (Head et al., this issue). Like the Quaternary, it has often been regarded as a full System/Period (Salvador, 2006a, b), but discussions are in progress on whether the Neogene and Paleogene should be reclassified as sub-systems/sub-periods.

The chart extends to 2.7 million years to include the very end of the preceding Piacenzian Stage of the Pliocene Series.

Pleistocene GSSPs

Formal GSSPs for the Pleistocene Subseries will be proposed in the near future. The INQUA Commission on Stratigraphy/ICS Working Group on Major Subdivision of the Pleistocene agreed to place the Early/Lower–Middle boundary at the Brunhes/Matuyama magnetic reversal Chron boundary (Richmond, 1996). A stratotype locality has yet to be identified, but three candidate sections are being considered by SQS Working Group (Head et al., this issue). Following recent re-evaluation, the Middle–Late/Upper boundary is placed, following historical precedent in NW Europe, at the Saalian–Eemian Stage boundary. The former is positioned at the basal-boundary stratotype of the Eemian in the Amsterdam–Terminal borehole, the Netherlands (Gibbard, 2003; Litt & Gibbard, this issue).

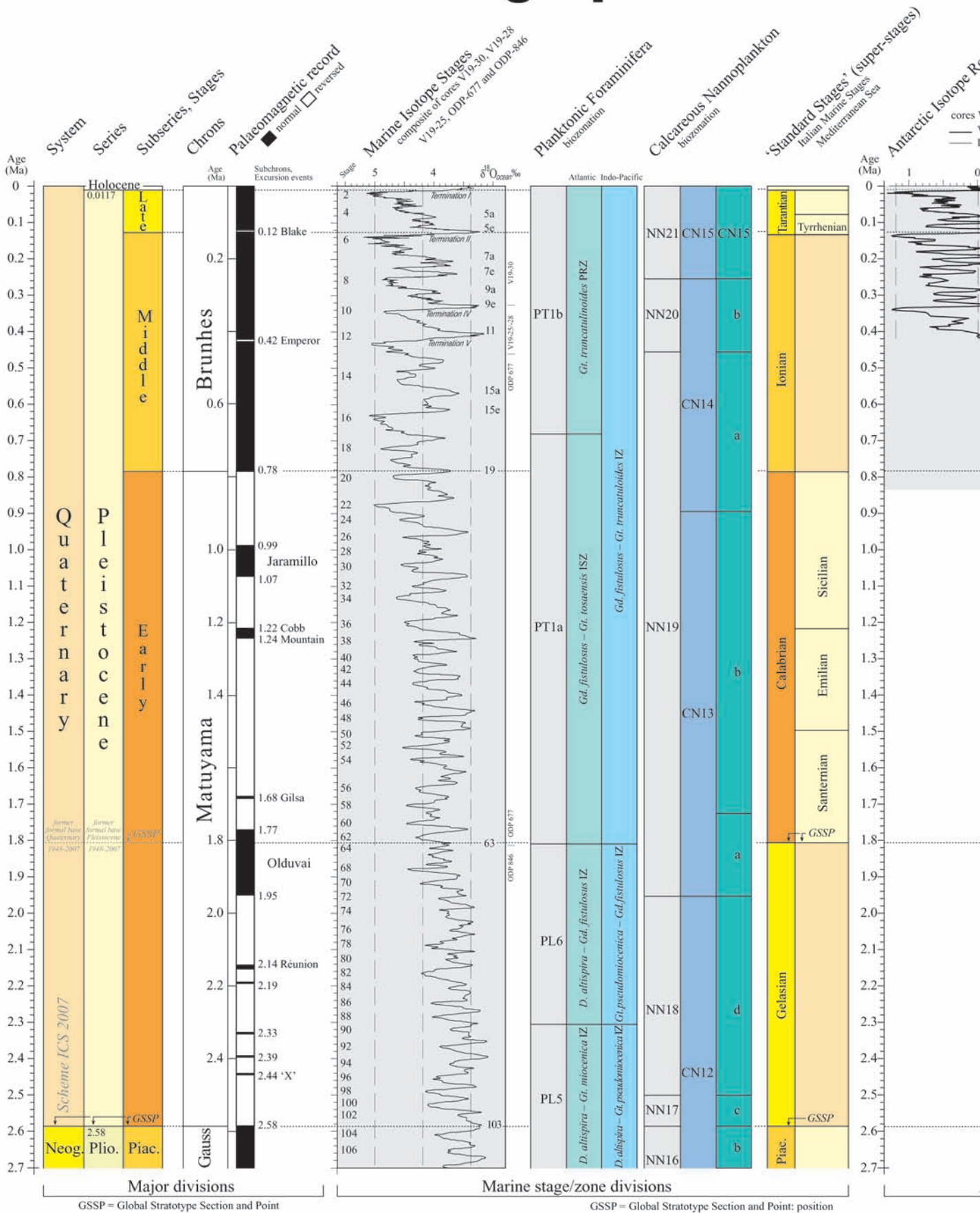
The Holocene is generally regarded as having begun 10,000 radiocarbon years before 1950 AD, or 11.7k calendar years before 2000 AD (Wolff, 2008). This boundary has been defined as a Global Stratotype Section and Point (GSSP; Walker et al., this issue) in the North-GRIP ice core of the Greenland Ice-Core Project (NGRIP; Rasmussen et al., 2006). Auxiliary stratotypes are also defined, for example, in an annually-laminated lake sequence in western Germany (Litt et al., 2001).

Marine stage/zone divisions

Isotope studies from the bottom sediments of the world's oceans have indicated that as many as 52 cold and interspersed warm climate periods, often referred to as glacials and interglacials, occurred during the last 2.6 million years. In contrast to the deep sea, continental evidence is so incomplete and regionally variable that terrestrial glacial–interglacial stratigraphies must refer to the ocean record for a global chronological foundation.

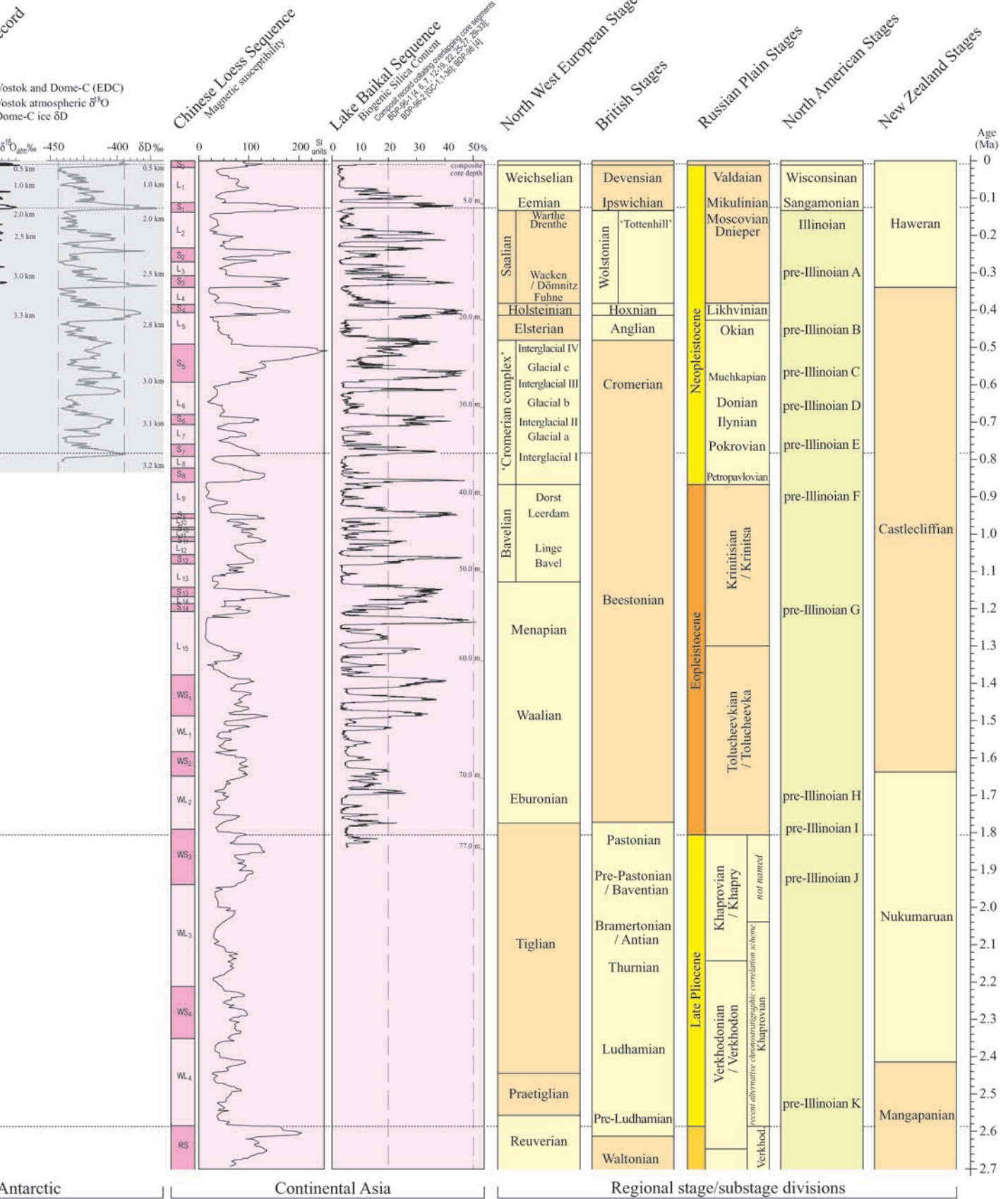
Here the deep-sea based, climatically-defined chronostratigraphy is taken from oxygen isotope data collected and processed by Crowhurst (2002), updated by Tzedakis et al. (2006). It is plotted against the magnetostratigraphic time scale prepared and modified from Funnell (1996). The curve plots depict $\delta^{18}\text{O}$ (the ratio of ^{18}O versus ^{16}O) in the tests of fossil benthonic (ocean-floor dwelling) foraminifera. Shifts in this ratio are a measure of global ice-volume,

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which is dependent on global temperature and which determines global sea-level. Planktonic foraminifera and calcareous nannoplankton provide an alternative biostratigraphical means of subdivision of marine sediments. The micropalaeontological zonation is taken from Berggren et al. (1995).

'Standard stage' ('super-stage') global divisions

The desire to divide Quaternary/Pleistocene time into 'standard stages', that is units of approximately the same duration as those in the pre-Quaternary Tertiary time, has been advocated on occasions. The only succession that has been divided in this way is the shallow marine sequences in the Mediterranean region, especially in southern Italy, based principally on faunal and protist biostratigraphy. For various reasons the scheme was considered unsatisfactory for use beyond this region. Renewed investigation in recent years has led to the proposal of units based on multidisciplinary investigation. The Italian shallow marine stages are derived from Van Couvering (1997) modified by Cita et al. (2006 and Cita, this issue).

In view of their duration, encompassing multiple climate cycles and periods for which regional stage units of markedly shorter duration have been defined, these 'standard stages' are considered as 'super-stages'.

Major continental records: Antarctic ice, Chinese Loess, Lake Baikal

Two plots of isotope measurements from Antarctic ice-cores are shown. The first is the 420 ka-long plot from the Vostok core and shows atmospheric $\delta^{18}\text{O}$ (Petit et al., 1999), determined from gas bubbles in the ice. This atmospheric $\delta^{18}\text{O}$ is inversely related to $\delta^{18}\text{O}$ measurements from seawater and therefore is a measure of ice-volume. It can also be used to separate ice volume and deepwater temperature effects in benthic foraminiferal $\delta^{18}\text{O}$ measurements. The deuterium measurements (δD) for the last 800 ka are from the 3.2 km deep EDC core in Dome C (EPICA community members, 2004; Jouzel et al., 2007). They come from samples of the ice itself and give a direct indication of Antarctic surface palaeotemperature.

For the Chinese loess deposits the chart shows the sequence of palaeosols (indicated by S and WS) and alternating loess units (L and WL) from Luochuan (An et al., 1990). It is accompanied by a continuous plot of magnetic susceptibility from the same sequence.

The Siberian Lake Baikal provides a bioproductivity record from the heart of the world's largest landmass, an area of extreme continental climate. High concentrations of biogenic silica indicate high aquatic production during interglacials (i.e., lake diatom blooms during ice-free summer seasons). The composite biogenic silica record from cores BDP-96-1, -96-2 and -98 is plotted on an astronomically tuned age scale (Prokopenko et al., 2006). The composite record extends well beyond the top of the Olduvai reversal, a tuned age scale for this part of the series is in preparation.

Regional stage/substage divisions

The continuous sequences, above, provide the comparison for a selection of continental and shallow marine stage-sequences from around the world reconstructed from discontinuous sediment successions. Solid horizontal lines on the plots indicate observed boundaries, where no lines separate stages, additional events may potentially be recognised in the future.

The NW European stages are taken from Zagwijn (1992) and De Jong (1988). The British stages are taken from Mitchell et al.

(1973); Gibbard et al. (1991) and Bowen (1999). The Russian Plain stages are taken from the Stratigraphy of the USSR: Quaternary System (1982, 1984), Krasnenkov et al. (1997), Shik et al. (2002), Iossifova (pers. comm.) and Tesakov (pers. comm). In addition, the Russian Pleistocene is also frequently divided into the Eopleistocene, equivalent to the Early Pleistocene Subseries, and the Neopleistocene, equivalent to the Middle and Late Pleistocene Subseries. The North American stages are taken from Richmond (unpublished). The New Zealand stages are from Pillans (1991) and Beu (2004).

Anthropocene

A recent proposal has been made to establish a new series status division following the Holocene, to be termed the *Anthropocene*. The term is being increasingly employed to identify the current interval of anthropogenic global environmental change, and may be adopted on stratigraphical grounds. It might be adopted at formal Series/Epoch level, for the time since the start of the Industrial Revolution when changes sufficient to leave a global stratigraphic signature distinct from that of the Holocene or of previous Pleistocene interglacial phases, occurred. A boundary definition may be made either using a Global Stratigraphic Section and Point ('golden spike') localities, and or by adopting a numerical date (Global Standard Stratigraphic Age: GSSA). Formal adoption of this term depends on its utility, particularly to Earth Scientists working on late Holocene successions (Zalasiewicz et al., 2008). Because of its short time duration, it has not been included in the chart.

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